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**Procedure for Probabilistic Tsunami  
Hazard Assessment for  
Incomplete and Uncertain Data**



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# Fundamental Question



How to assess the likelihood of such an event??





# Proposed Approach

- ❖ Empirical approach based on tsunami **observations** which are often **incomplete** and **uncertain**.
- ❖ Alternative to empirical approach is **tsunami propagation modelling**.  
(Comprehensive review of alternative approaches e.g. Geist, 2009)

# Problems Faced Using Only Observations



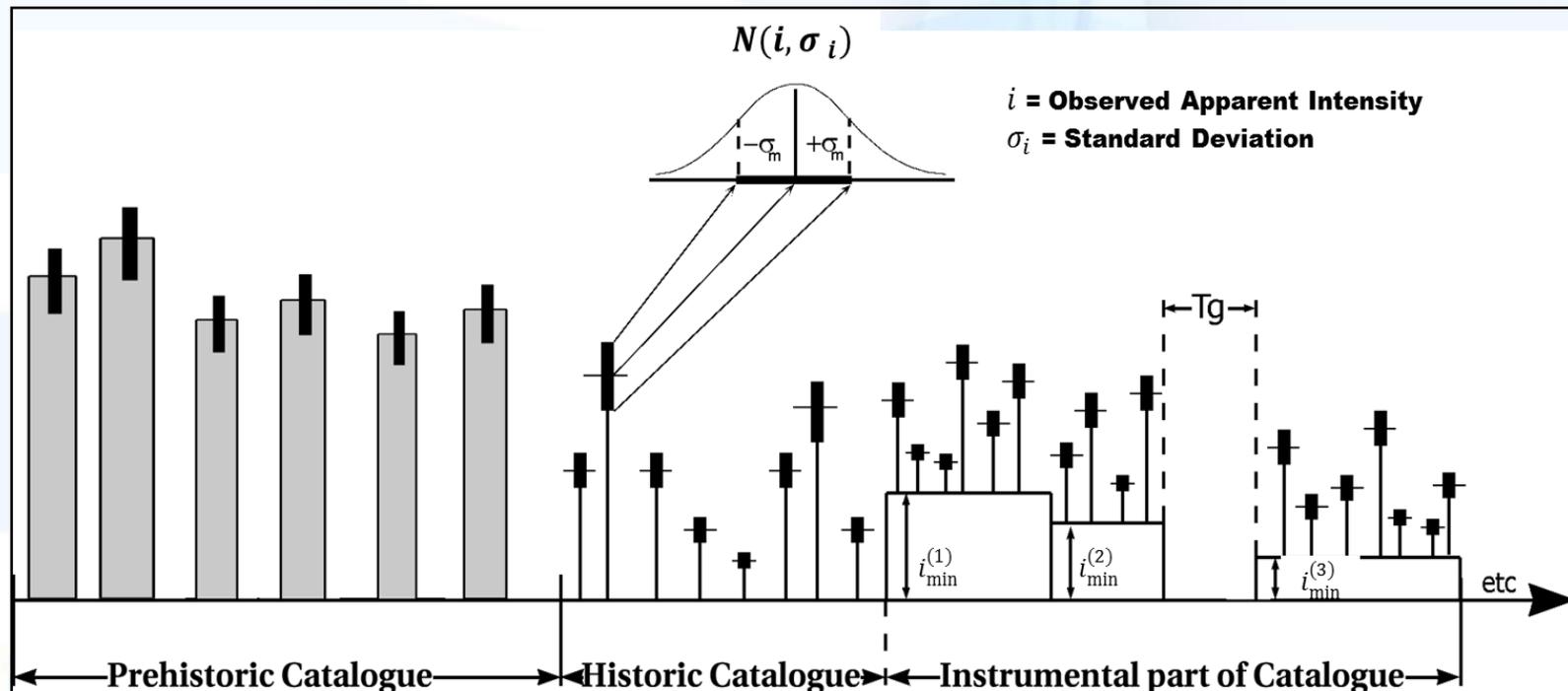
1. **Incompleteness** of tsunami database.
2. Uncertainty in the **intensity** (magnitude/size) of an event.
3. Uncertainty in the applied **occurrence model**.
4. **Maximum** local/regional tsunami intensity.





# Nature of Tsunami Input Data

- ❖ Very strong pre-historic events (palaeo) – last thousands of years.
- ❖ Strongest historic events – last few hundred years.
- ❖ Recent events – in last 100 years (complete catalogue).





# Model Parameters

- ❖ Tsunami hazard described by probabilities and return periods is characterised by 3 parameters:

$\lambda$	Mean activity rate of tsunami occurrence
$b$ -value	Ratio between strong and weak intensity tsunamis (equivalent to the frequency-magnitude G-R relation)
$i_{\max}$	Upper limit of tsunami intensity (or tsunami run-up)





# Proposed PTHA Model

The proposed methodology addresses:

A. Temporal distribution.

B. Intensity distribution.

❖ Discrepancy between data and occurrence model.

C. Combination of Extreme and Complete catalogues.





## A. Temporal Distribution of Tsunamis

Prob [ $n$  tsunami events, observed in time interval  $t$  along a coast]

$$p(n|\lambda, t) = \frac{(\lambda t)^n}{n!} \exp(-\lambda t) \quad n = 0, 1, \dots$$



Poisson-Gamma **PDF**

$$p(n|\lambda, t) = \frac{\Gamma(n + q_\lambda)}{n! \Gamma(q_\lambda)} \left( \frac{p_\lambda}{t + p_\lambda} \right)^{q_\lambda} \left( \frac{t}{t + p_\lambda} \right)^n$$

where

- $p_\lambda = \bar{\lambda} / \sigma_\lambda^2$
- $q_\lambda = \bar{\lambda} p_\lambda$



## B. Distribution of Tsunami Intensity

### Assume

- Observed intensity  $i$  = unknown “true” intensity + random error
- Random error can be significant  $\sim$  Gaussian (0, SD)
- Intensity of tsunami occurrence often very uncertain

(Tinti & Mulargia, 1985)

**...especially for palaeo and historic events**



## B. Distribution of Tsunami Intensity

### Assume

Tsunami run-up heights ( $h$  in meters) follows negative exponential distribution similar to G-R relationship

- ❖ Soloviev (1969) introduced a frequency-size distribution

$$\lambda = a \times 10^{-bi}$$

were  $a$  and  $b$  are regression coefficients.

- ❖ Intensity  $i$  can be linked with tsunami run-up ( $h$ ) along a coastline

$$i = \log_2(\sqrt{2h})$$



## B. Distribution of Tsunami Intensity

**CDF** of tsunami intensity

$$f_I(i) = \begin{cases} 0 & i_{\min} < i \\ \frac{\beta \exp[-\beta(i - i_{\min})]}{1 - \exp[-\beta(i_{\max} - i_{\min})]} & \text{for } i_{\min} \leq i \leq i_{\max} \\ 0 & i > i_{\max} \end{cases}$$



Compound **CDF**

$$f_I(i|i_{\min}) = \left[ 1 - \left( \frac{q_\beta}{q_\beta + \bar{\beta}(i_{\max} - i_{\min})} \right)^{q_\beta} \right]^{-1} \left[ \bar{\beta} \left( \frac{q_\beta}{q_\beta + \bar{\beta}(i - i_{\min})} \right)^{q_\beta + 1} \right]$$

where  $q_\beta = \bar{\beta} p_\beta$





## C. Combination of Catalogues

For each part of tsunami catalogues build likelihood function:

- $L_{\text{Paleo}}(\boldsymbol{\theta})$  = likelihood function based on **palaeo** tsunamis
- $L_{\text{Historic}}(\boldsymbol{\theta})$  = likelihood function based on **historic** tsunamis
- $L_{\text{Complete}}(\boldsymbol{\theta})$  = likelihood function based on **complete** tsunami catalogues

**Total likelihood function:**

$$L(\boldsymbol{\theta}) = L_{\text{Paleo}}(\boldsymbol{\theta}) \times L_{\text{Historic}}(\boldsymbol{\theta}) \times L_{\text{Complete}}(\boldsymbol{\theta})$$



## Parameter Estimation of $\lambda$ and $\beta$

$$L(\boldsymbol{\theta}) = L_{\text{Paleo}}(\boldsymbol{\theta}) \times L_{\text{Historic}}(\boldsymbol{\theta}) \times L_{\text{Complete}}(\boldsymbol{\theta})$$

Obtain parameter estimates  $\bar{\lambda}$  and  $\bar{\beta}$  through the

### *Maximum Likelihood Estimation Method*

which for given  $i_{\text{max}}$  maximizes the likelihood function  $L(\boldsymbol{\theta})$ .

Obtained by solving:

$$\begin{bmatrix} \frac{\partial \ln[L(\boldsymbol{\theta})]}{\partial \bar{\lambda}} \\ \frac{\partial \ln[L(\boldsymbol{\theta})]}{\partial \bar{\beta}} \end{bmatrix} = 0 \quad \text{where } \boldsymbol{\theta} = (\bar{\lambda}, \bar{\beta}, i_{\text{max}})$$



## Parameter Estimation of $i_{\max}$

$$i_{\max} = i_{\max}^{obs} + \Delta$$

where  $\Delta = \frac{\delta^{1/q_\beta} \exp[nr^{q_\beta}/(1-r^{q_\beta})]}{\bar{\beta}} \left[ \Gamma\left(-\frac{1}{q_\beta}, \delta r^{q_\beta}\right) - \Gamma\left(-\frac{1}{q_\beta}, \delta\right) \right]$

(Kijko, 2004; Kijko and Singh, 2011)

- $r = \frac{p_\beta}{p_\beta + i_{\max} - i_{\min}}$

- $C_\beta = \left[ 1 - \left( \frac{q_\beta}{q_\beta + \bar{\beta}(i_{\max} - i_{\min})} \right)^{q_\beta} \right]^{-1}$

- $\delta = nC_\beta$

- $p_\beta = \bar{\beta}/\sigma_\beta^2 \quad \& \quad q_\beta = (\bar{\beta}/\sigma_\beta)^2$

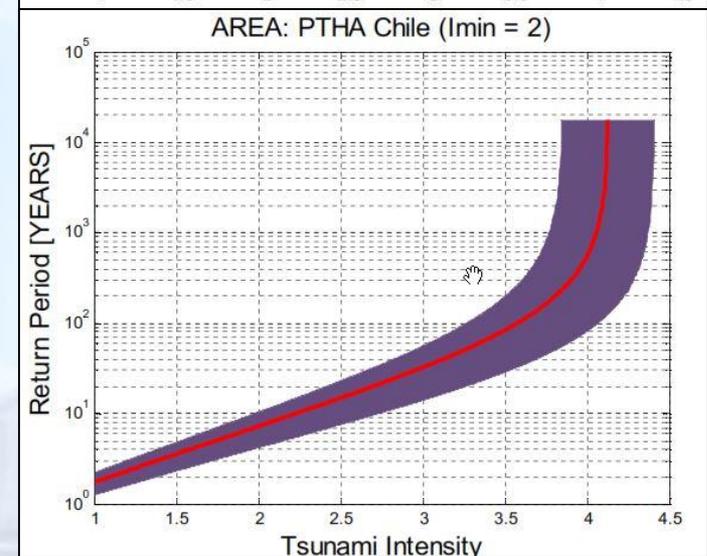
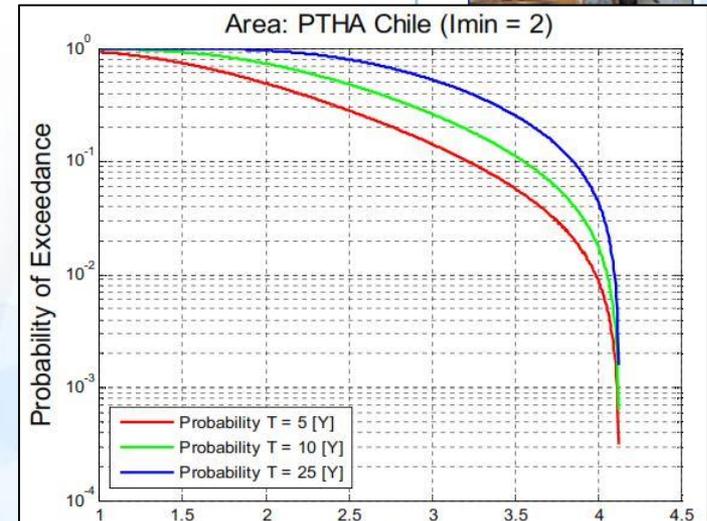
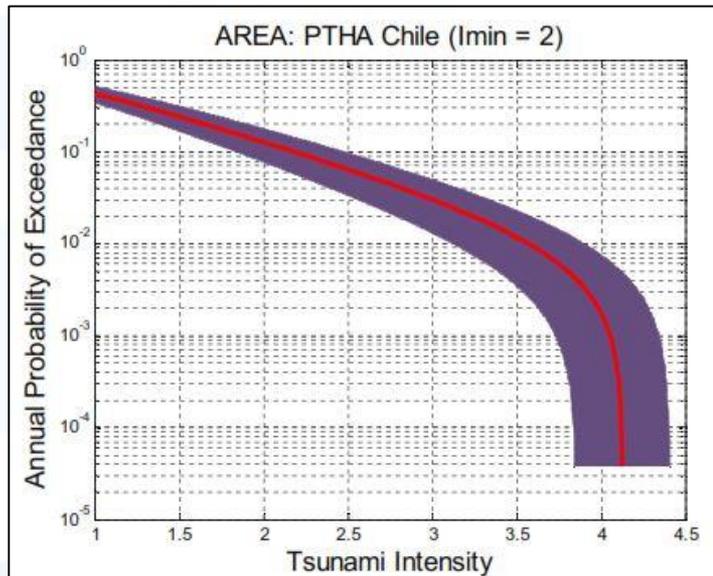
$\Gamma(\cdot, \cdot)$  = complementary Incomplete Gamma Function (Abramowitz and Stegun, 1970)





# Example 1: Chil 

Only Extreme (Historic) events catalogues



Tsunami database of Dr V.K. Gusiakov,  
Institute of Computational Mathematics and  
Geophysics, Siberian Division, Russian Academy  
of Sciences, Russia

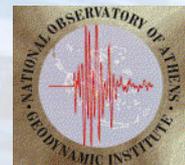


## Example 2: Areas surrounding Greece

Extreme (historic) and complete events catalogues for

- Hellenic Arc
- Corinth Gulf
- Mediterranean Sea

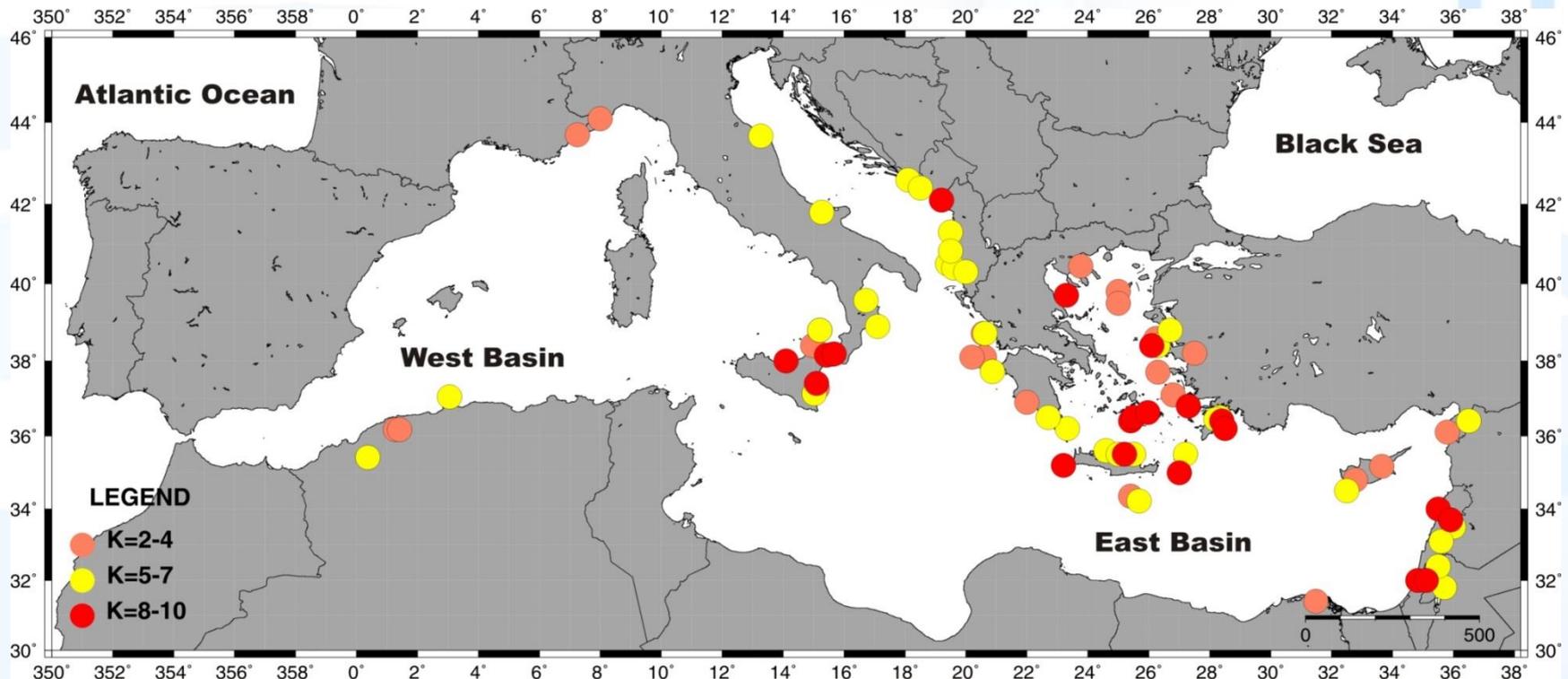
Tsunami database of Prof G.A. Papadopoulos,  
Institute of Geodynamics, National Observatory of Athens, Greece.



# Applied Tsunami Catalogue



1620 BC – present – Mediterranean Sea



Tsunami database of Prof G.A. Papadopoulos, Institute of Geodynamics, National Observatory of Athens

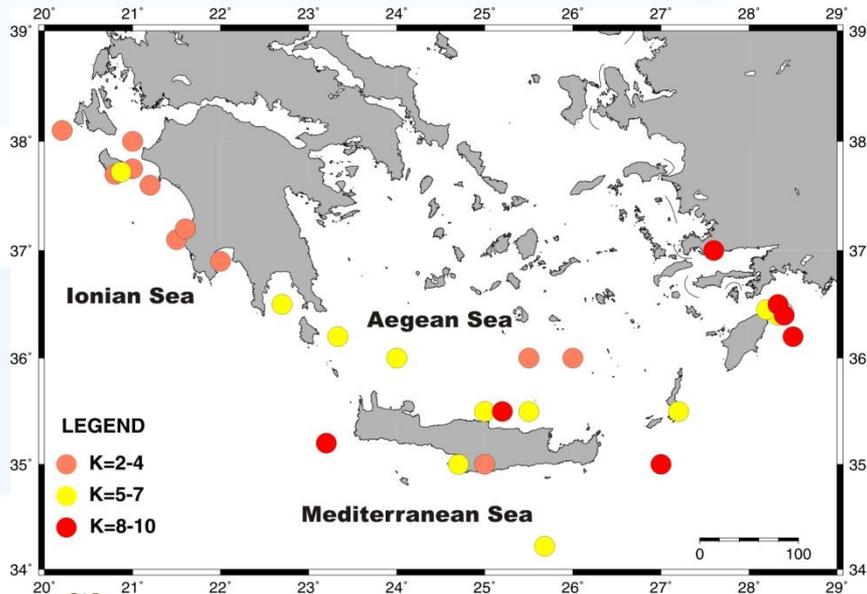


# Applied Tsunami Catalogue

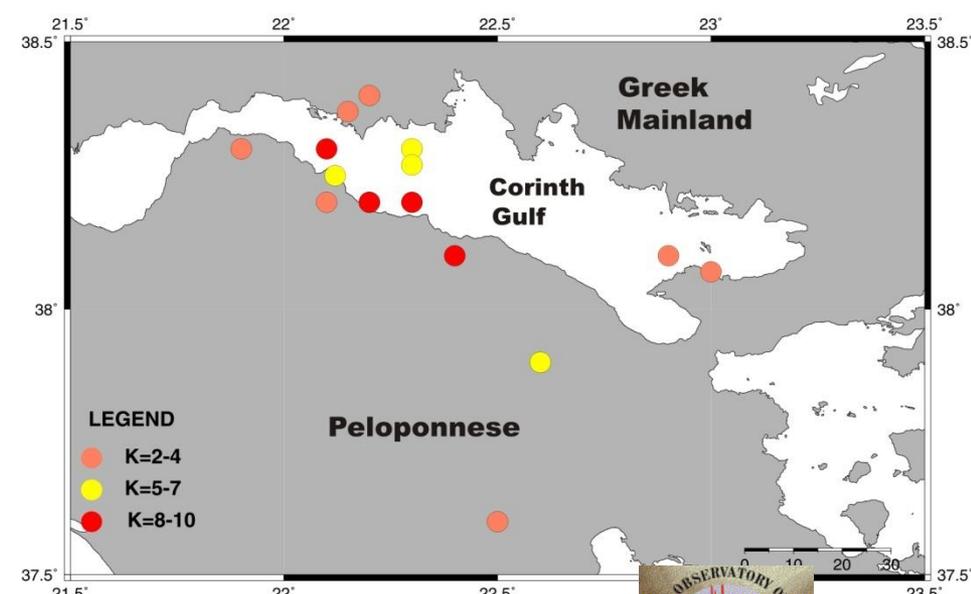
1620 BC – present – Hellenic Arc

373 BC – present – Corinth Gulf

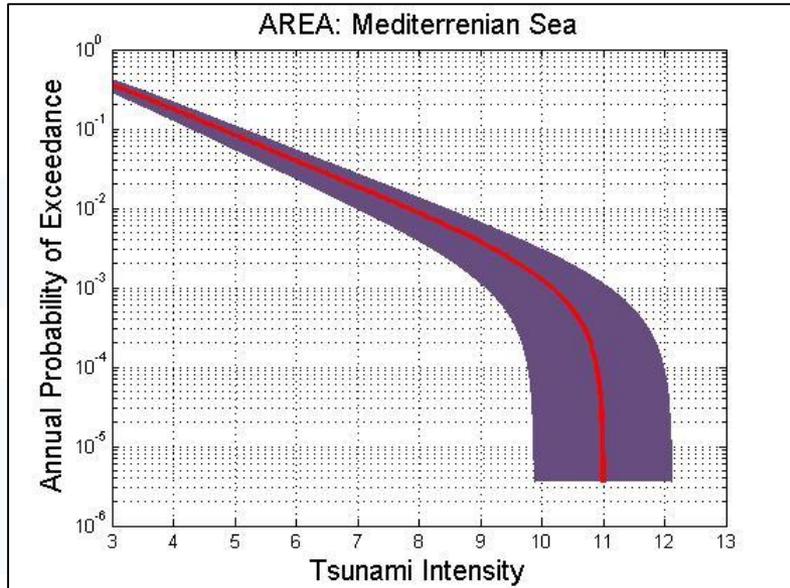
## Hellenic Arc



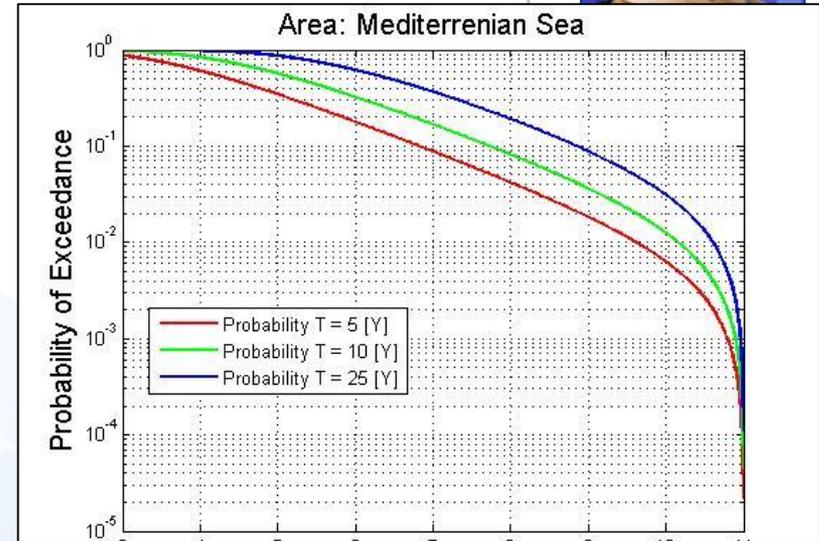
## Corinth Gulf



# Results: Mediterranean Sea

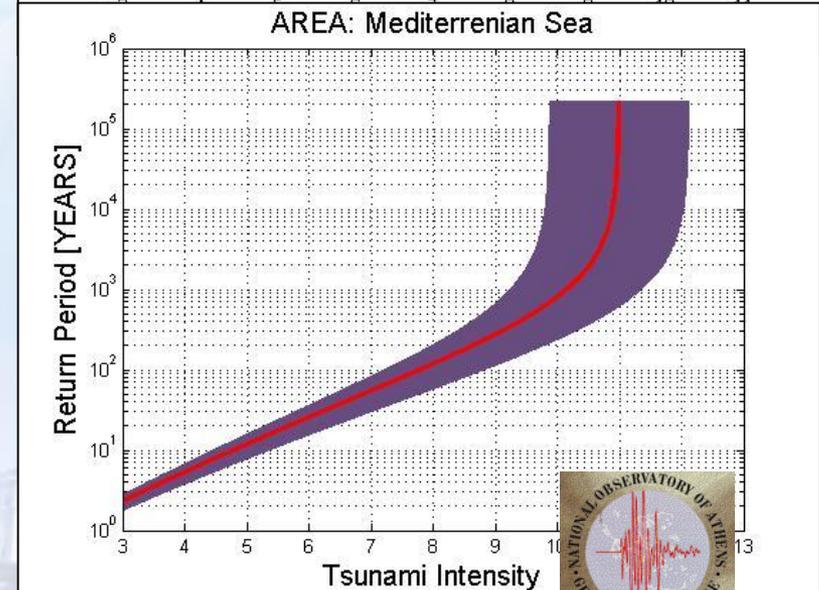


(a)



(b)

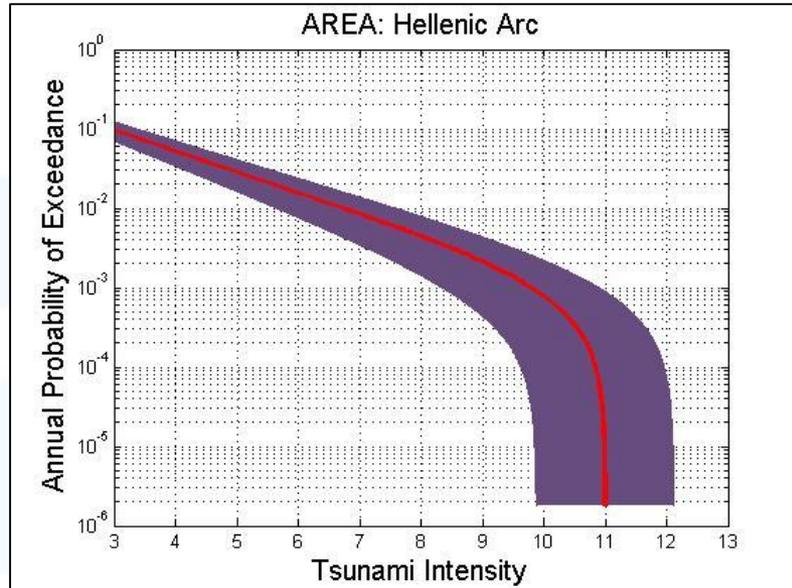
- a. Annual probability of exceedance
- b. Probability of exceedance for 5, 10 and 25 yrs
- c. Mean return period



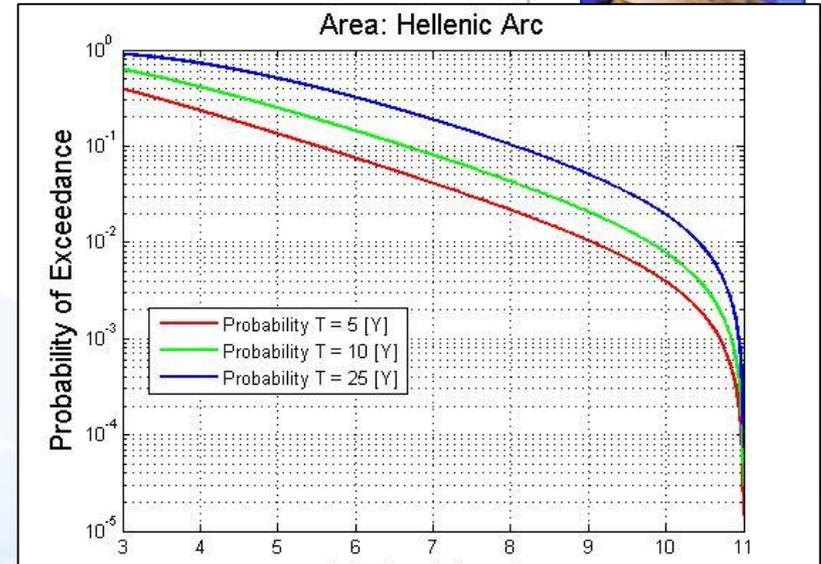
(c)



# Results: Hellenic Arc

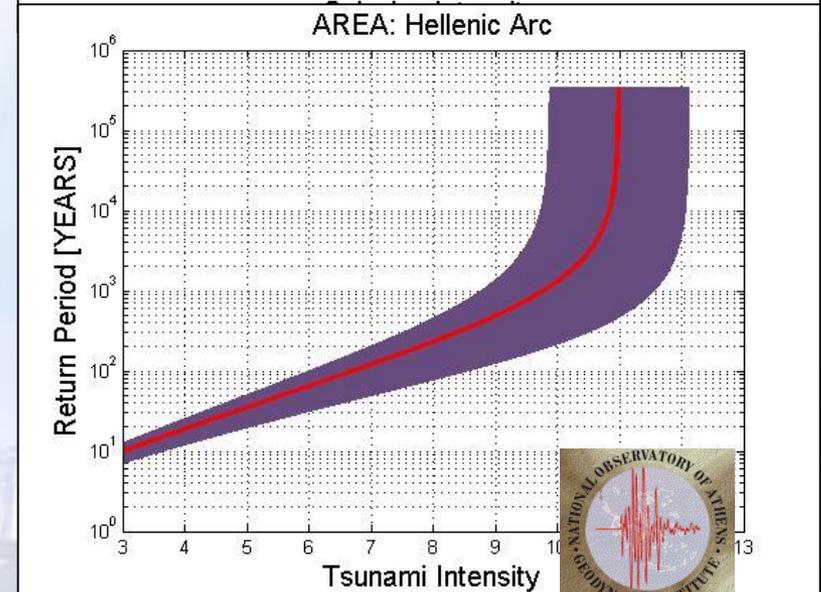


(a)



(b)

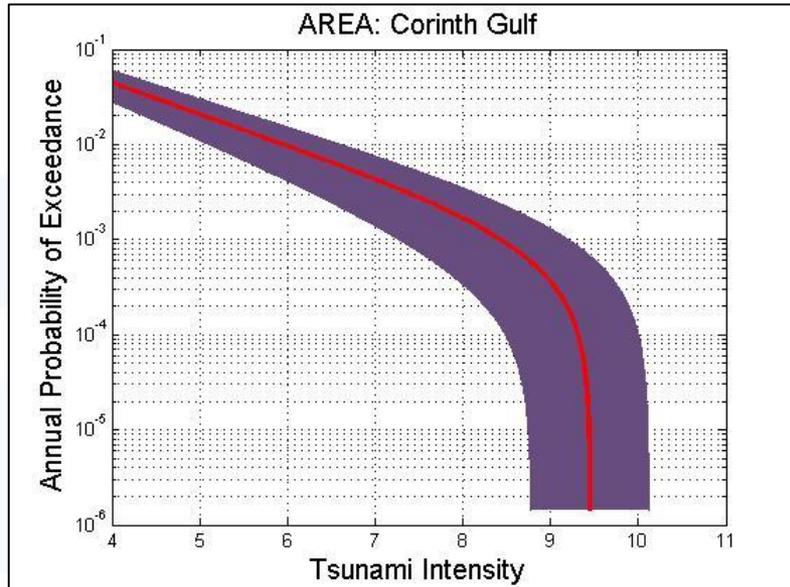
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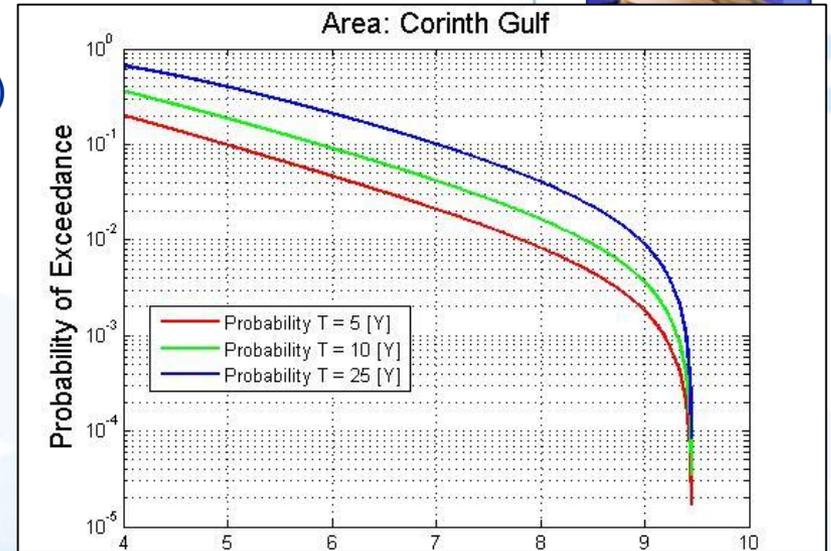
(c)



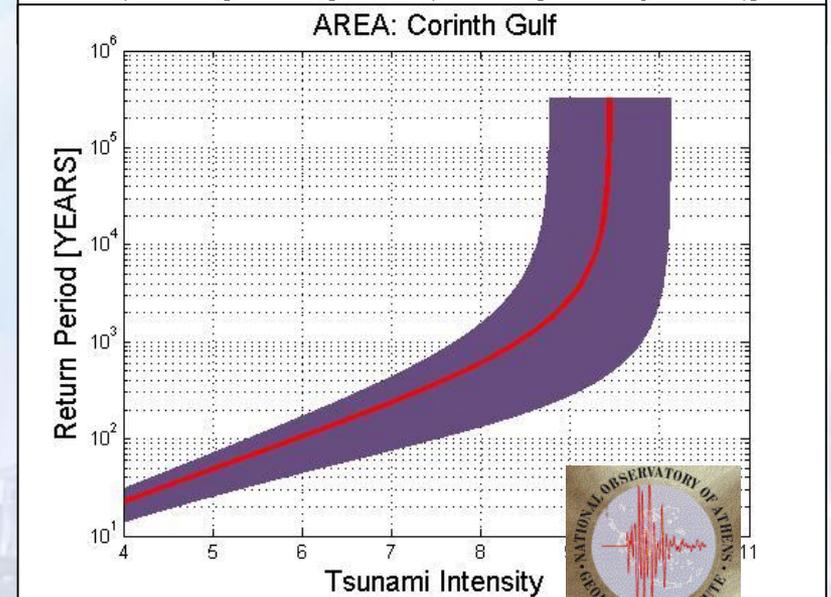
# Results: Corinth Gulf



(a)



(b)



(c)

- a. Annual probability of exceedance
- b. Probability of exceedance for 5, 10 and 25 yrs
- c. Mean return period



# Mean Return Periods

Return Period	Intensity 5.0	Intensity 9.0
Mediterranean Sea	11.7 yrs	271 yrs
Hellenic Arc	34.6 yrs	477 yrs
Corinth Gulf	48.2 yrs	2800 yrs



# Conclusions

- A new procedure for PTHA has been developed to calculate the tsunami hazard for a specified region that caters for **incomplete** and **uncertain** data.
- The procedure:
  - ❖ permits the assessment of the **maximum likelihood** estimates of the key tsunami hazard parameters.
  - ❖ is **flexible** and allows for the use of palaeo, historic and complete tsunami catalogues.
- Do not need an extensive and complete events catalogue





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